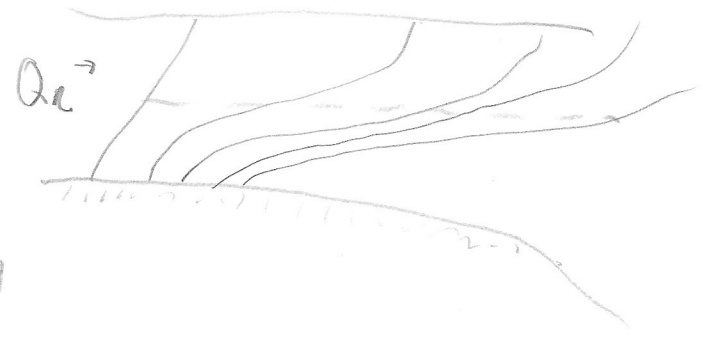
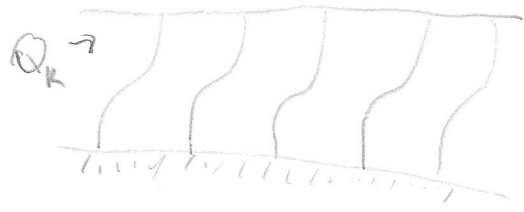


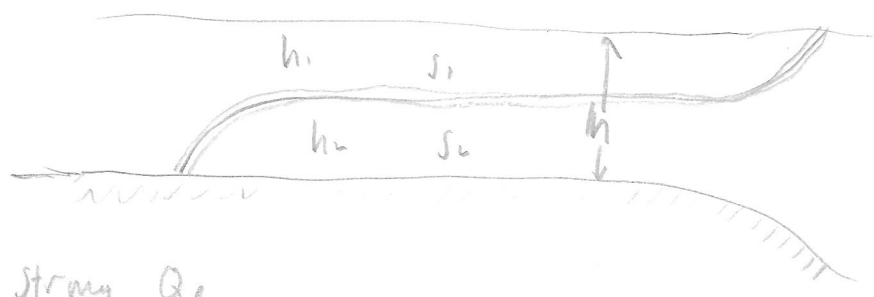
RG Strongly stratified "two-layer" estuaries

8/8/2019

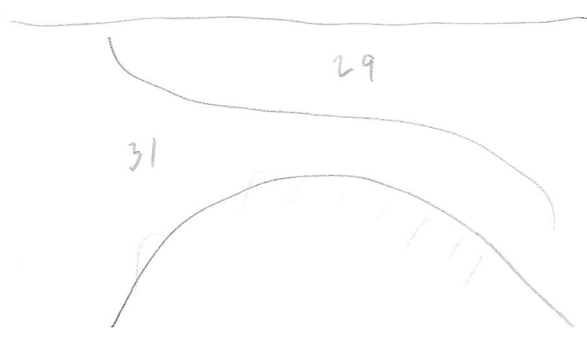
(1)



as it gets more stratified
 the "identity" of the water depends on
 which layer it is in

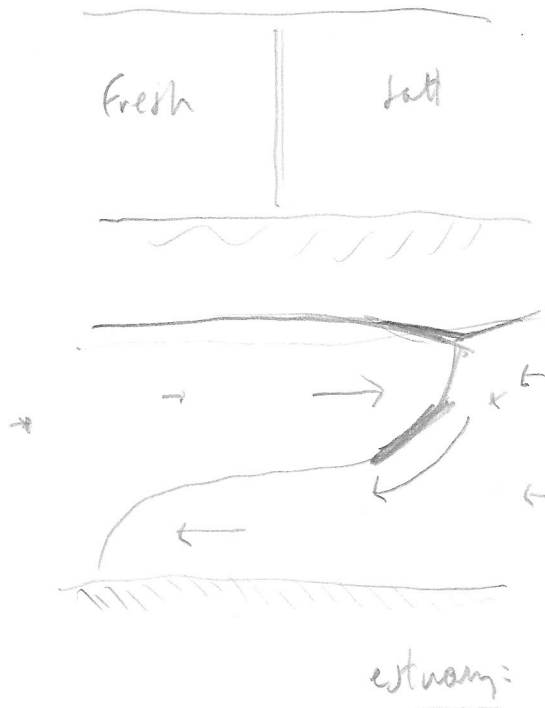


- Strong Q_e
- Weak mixing
- water does not stay in the estuary very long
 so it doesn't have time to mix

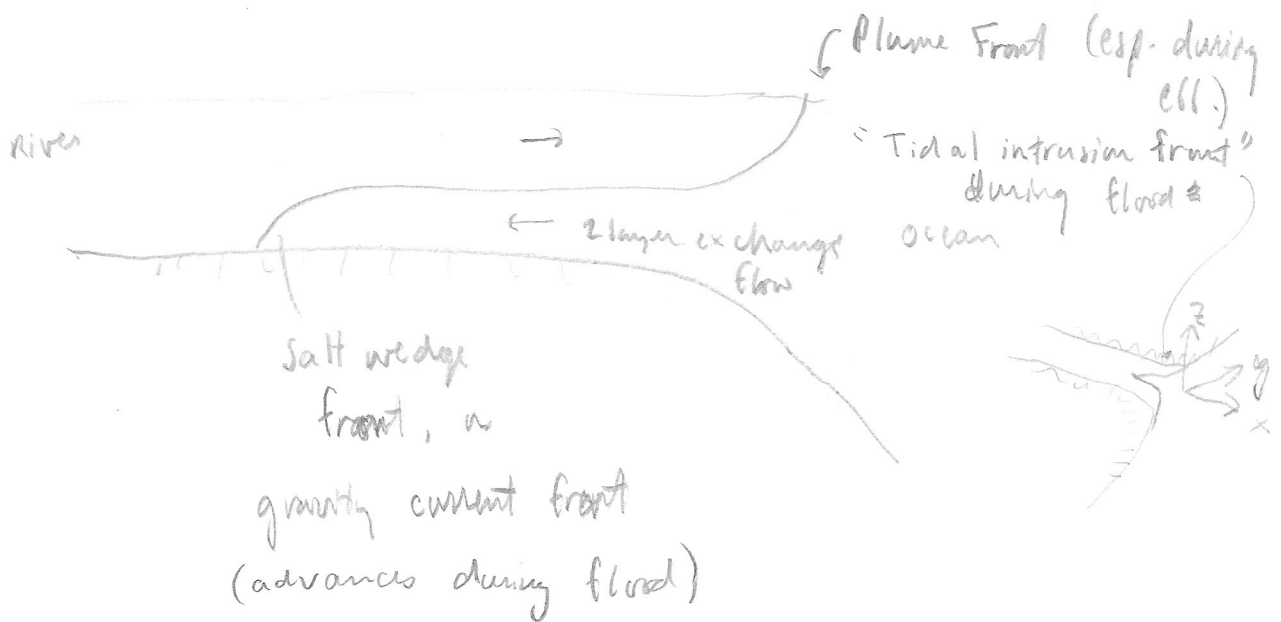


Flow over a sill
 - 2 layers were through
 1 is not fresh.

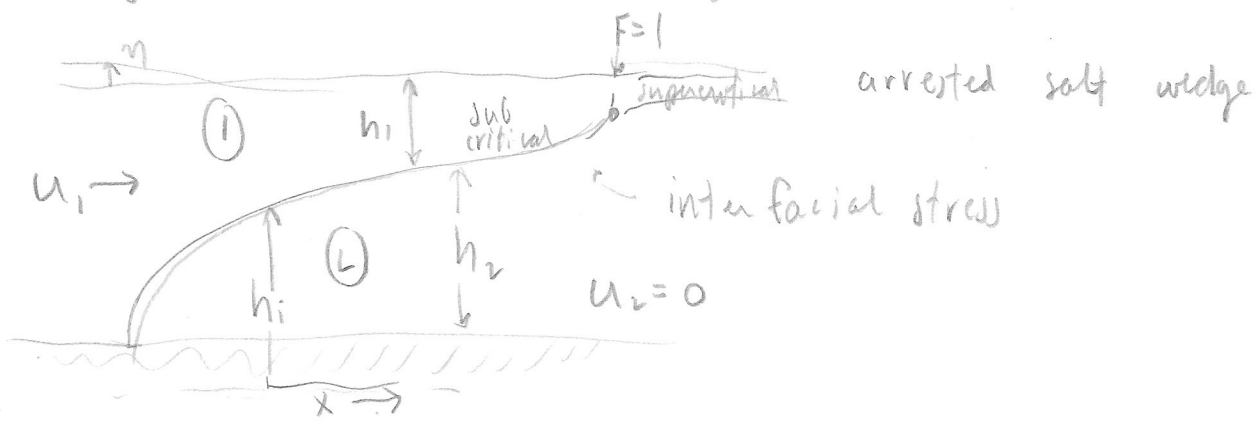
lock exchange problem



the two types of water may have been created elsewhere, and then meet because of some physical situation, like flow separation.



Schijf & Shoemaker (1953) solutions



$$\Delta u = u_1 - u_2$$

$$u_1 u_{1x} + g \eta_x + \frac{C_i |\Delta u| \Delta u}{h_1} = 0$$

$$g \eta_x + g' h_{ix} - \frac{C_i |\Delta u| \Delta u}{h_2} = 0$$

Interfacial friction but no entrainment (actually, ~~what~~ entrainment there is in the interface - and is removed rapidly?)

$h_{ix} \sim 50$ times steeper than η_x

$$C_i \approx 1-7 \times 10^{-4}$$

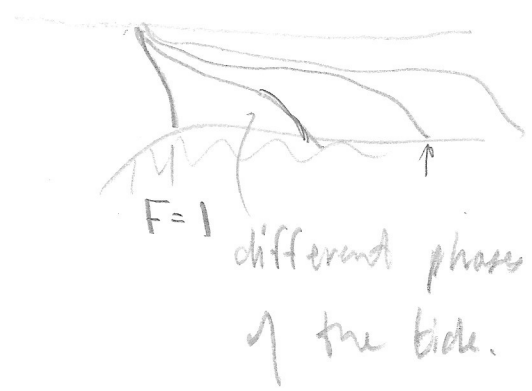
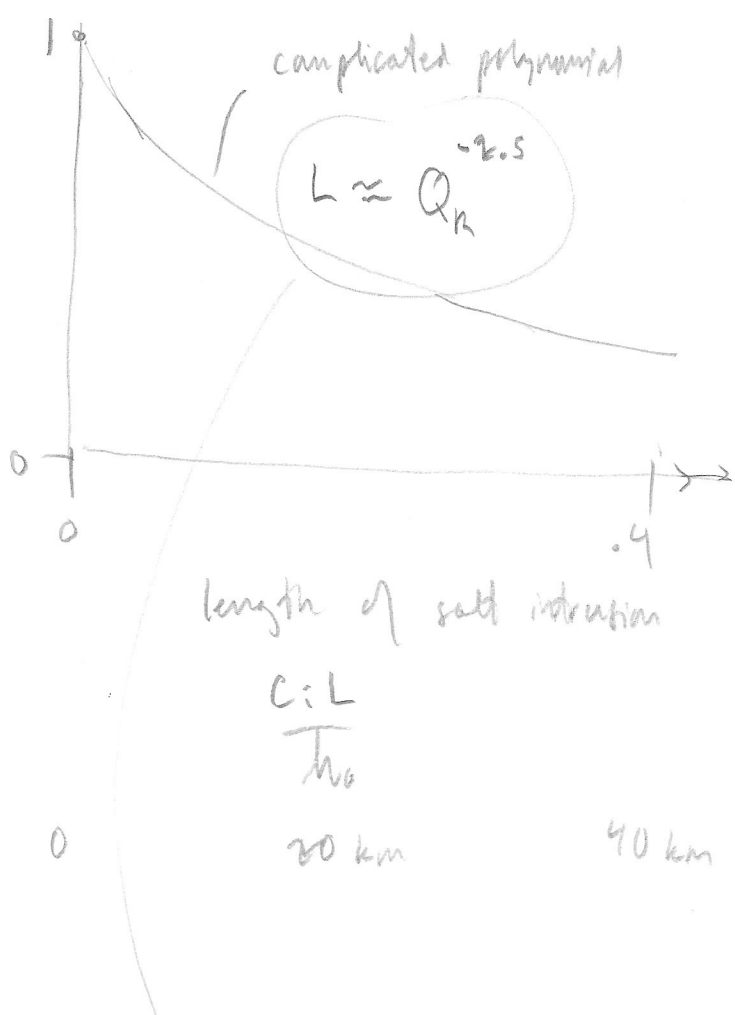
Froude # $Fr_f = \frac{u_f}{\sqrt{g' h_0}} \sim 1$

$$u_f = u_r + u_t$$

$$\left\{ \begin{array}{l} u_r \sim 0.4 \text{ m/s} \\ u_t(\text{tidal}) \sim 1 \text{ m/s} \end{array} \right\} 1.4 \text{ m/s}$$

different answers at other phases of the tide

$$F = \frac{U_e}{(g'h_0)^{1/2}}$$



so this is much more sensitive to Q_R compared to the partially mixed result $L \sim Q_R^{-1/3}$

It is more sensitive because of the advective terms.

$$u_t + u_1 u_{1x} + g \eta_x \dots$$

$$\omega [u_1] \quad \frac{[u_1^2]}{L}$$

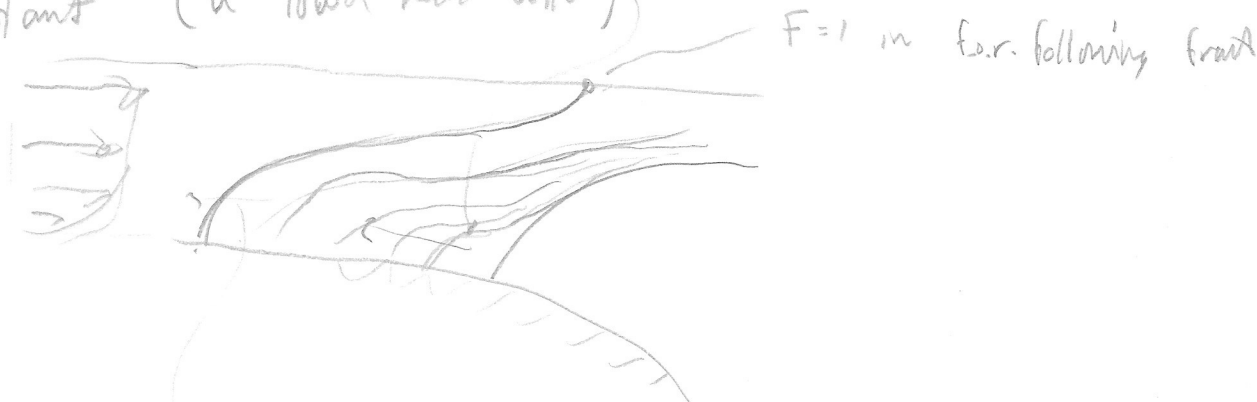
$$\frac{(a)}{(b)} = \frac{\omega L}{u} \sim 1.4$$

$1.4 \times 10^{-9} \text{ s}^{-1}$
 $10^4 \text{ m (Fram River)}$
 1 m/s

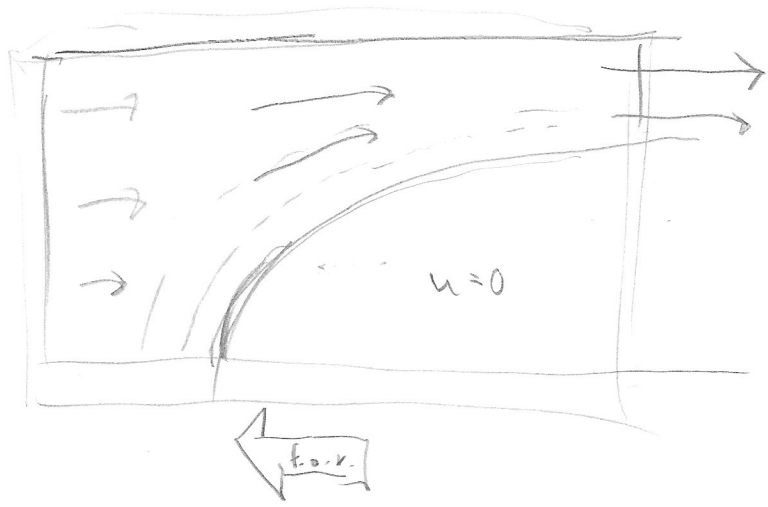
\Rightarrow at the peak tidal acceleration time dependence is important



Advancing salt wedge - the fact that there is a lag layer in the river is important (a lower near bottom)



Flood: ~ parabolic shape
 move up estuary without changing shape



In f.o.r.
moving with
the intruding
salt wedge front

① $u_1 u_{1x} + \frac{g \eta}{2} = 0$ neglected for thick upper layer
 $\Rightarrow 0$ so u_1 small

② $g' h_{1x} + \frac{C_0 u_2 |u_2|}{h_2} = 0$

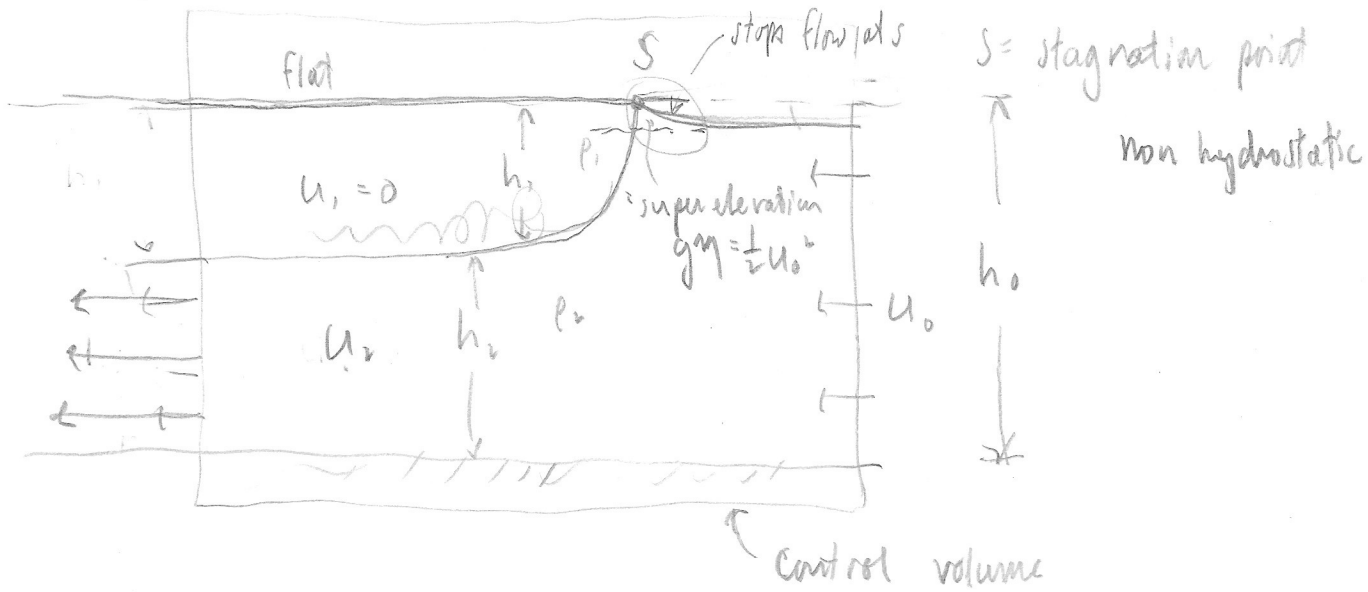
$\approx g' h_{2x} + \frac{C_0 u_2 |u_2|}{h_2} = 0$

at the nose the interface gets steeper
 as the friction term increases like $\frac{1}{h_2}$

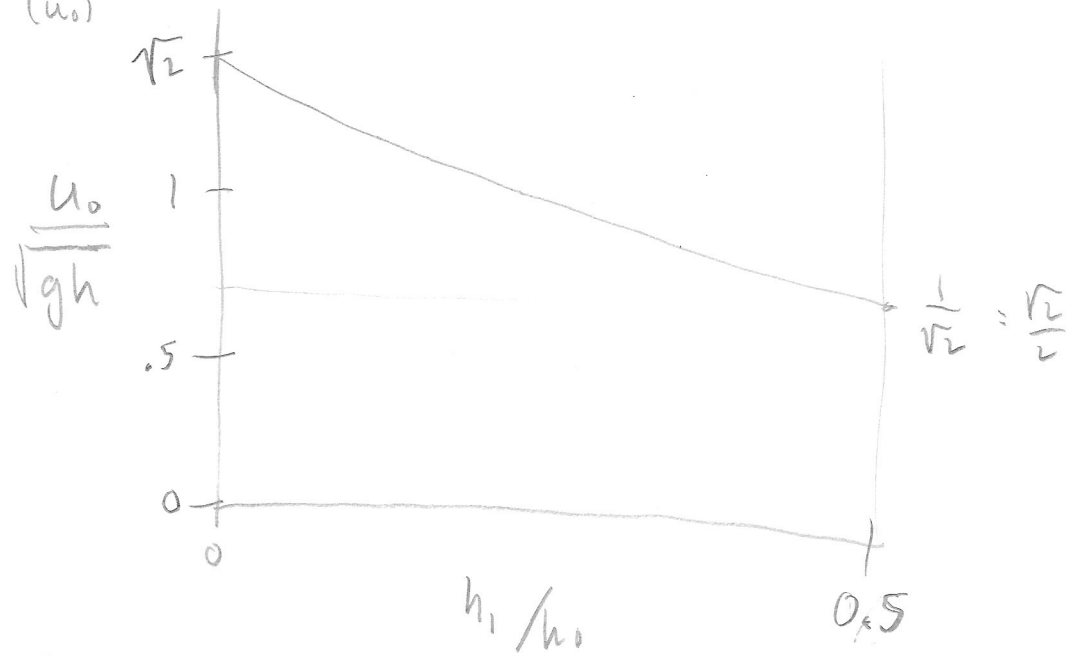
$h_2 \sim \sqrt{x}$

Nose is non-hydrostatic

Benjamin relation to gravity current (in far. of current nose)



speed of the front (u_0)

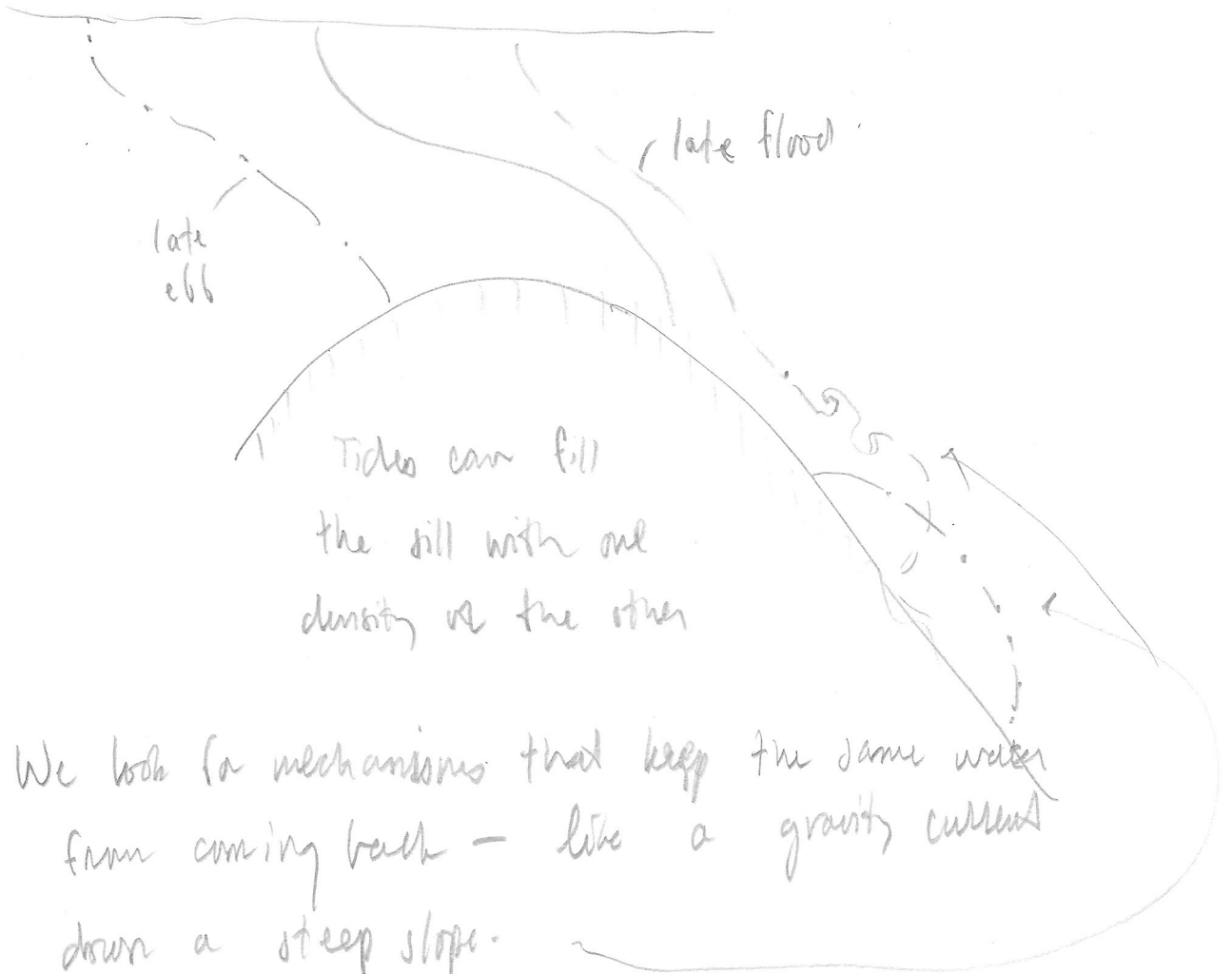


fastest flow has

$\frac{h_1}{h_0} \sim \frac{1}{3}$ and has greatest energy loss

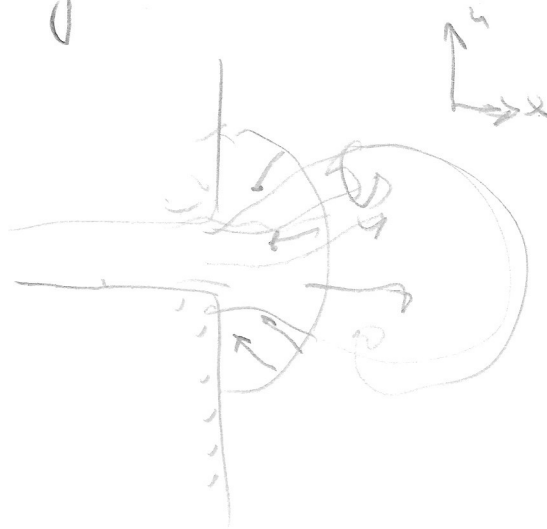
Fjord w/ tides

8



We look for mechanisms that keep the same water from coming back - like a gravity current down a steep slope.

Or plume dynamics



Benjamin → Website